CHEMICAL MIGRATION IN SNOWPACK

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INTRODUCTION:

to space and time; this variation may be even greater where orographic influences are present. It is inviting to use snowpack sampling as a technique to collect precipitation speci- while several periods of melting occurred at mens, and to evaluate chemical precipitation theories or source-receptor pollution transsis. Such snowpack sampling would allow a pos-spect to water was nearly conserved, when the for analysis, rather than requiring multi-point and Michigan, there was considerable exchange teriori collection of representative samples multi-time collections by several observers, through a long precipitation period.

Warburton and Linkletter [1978], measured chemical profiles in snowpits on the Ross Ice Shelf, Antarctica, and were able to identify major storm systems, from the Na, K and Mg pro- istry as a function of cloud physics. An exfiles in the snow. Their analysis indicated that riming processes near the coast greatly influenced deposit of both water and sea salt particles, and they inferred exchange and mixing processes on the interior of the shelf. Gjessing [1977, 1984, 1989] examined the chemical composition of polar snow in the Arctic and Antarctic. He accounted for a deficit of SO4 in coastal Antarctic snow through riming processes. He was able to relate chemical dep osition in Arctic snowfields to atmospheric trajectories. Twickler, et al. [1986] found a seasonal deposition pattern of organic carbon in Greenland snow samples collected near DYE3, Greenland (65.01°N, 44.87°W), with higher concentrations observed in the winter/spring period. Schaefer[personal communication] investi-optical properties, and several gas concentragated newly fallen snows, and found that the crystal type was preserved for as long as 24 hours allowing application of these snow crystals as aerological sondes.

Hogan, et al. [1985] attempted to synthesize chemical precipitation from analysis of new fallen snow and snowpack. This work differed from that of Cadle, et al. [1984], as new fallen snow was partitioned as a function

of meteorological conditions at one observation point, and snow pit collections made in Chemical precipitation varies with respect terms of discrete storm strata, rather than by length of core. Additionally, the air temperature remained below freezing at Whiteface through the accumulation of the pack, the Michigan site.

> Although chemical concentration with reentire pack is considered at both Whiteface of material among strata (or core segments).

> This apparent exchange within strata representing a single storm event, or among several strata, confounds any attempt at a posteriori examination of precipitation chemperiment has been initiated to investigate chemical behavior in snowpack.

INSTRUMENTATION:

Two permanent observatories are maintained at Whiteface Mountain; additional observation and sampling sites have been operated at several levels, seasonally, or in some cases, for a few years. The Schaefer Observatory at the mountain summit (4867 ft ASL, approximately the 850 mb level) is instrumented for cloud physics and cloud chemistry research and is frequently manned by observers, in addition to the recording instruments. Meteorological parameters, aerosol concentration, along with tions have been continuously measured for several years. Cloud parameters, and cloud water chemistry are measured when the observatory is immersed in cloud.

The observatory offices and laboratory analytical facilities are located at 1950 ft ASL on the east facing flank of the mountain. Several precipitation collectors, gas, and aerosol samplers (supported by DEC, ENSR and

these buildings. Chemical concentration in precipitation and summit cloud water are mea- pack in a covered watch glass on 25 February. sured soon after collection at this laboratory; Snowfall covered the specimen, and 3-10 cm additional collections are forwarded for anal- of snow remained atop the specimen from 4 ysis to the several national precipitation chemistry monitoring programs. The capabilities of the Whiteface observatories are tabulated in Table 1.

This suite of instruments, collectors, and analytical facilities provides an excellent setting for investigating the variation of concentration of components of precipitation from the cloud, into precipitation, and finally in residual snowpack. An instrumented snow field has been established 30M above the precipitation variables. A similar experiment was conducted in February and March of 1986 in an attempt to monitor snow stratigraphy and temof ions in a seasonal snowpack [Arons, personal hensive cloud-precipitation-snowpack chemical perature to follow changes in the distribution communication].

A rectangular grid (39m x 18m) was laid out above the precipitation collector area in early winter 1988-89. Four thermistor arrays were placed at the corners. The thermistor arrays sensed temperatures -3, 0, 3, 10, 30 and 60 cm relative heights from the snow-groundshallow snowpack. interface. The vertical temperature profile is used to determine the instantaneous absolute temperature, and the direction of heat flux within the snowpack. Two radiometers, one oriented vertically, and the other aligned withUSDA Forestry Service (Grant 23338). Gratithe slope, measure incoming direct normal and diffuse-sky radiation.

The thermistor and spectral pyranometer outputs are recorded on a Campbell 21X Scientific Data Logger. This data is transferred to personal computer for analysis. Examples of the temperature profiles recorded in the extremely sparse snowpack of 1989 are given in Figures 1, 2, and 3.

Additional to the recording instruments, four bulk precipitation storage collectors are gation of chemical conservation in a constrain-Dionne, B.C., Rounbehler, D.P., Achter, E.K., ed system approximating the environment of the pack.

RESULTS OF EXPERIMENTS:

The temperature histories shown in Figures 1, 2, and 3 indicate the dearth of snowpack maintained at Whiteface in 1989. The shallow snowpack and frequent thaws frustrated accomplishment of the primary experiment.

A tracer experiment was attempted using an organic solid (TNT) of vapor pressure

Log P (ppbv) = -5481/T + 19.37

MAPS3S) are situated in a flat area just above [Dionne, et al., 1986]. A few small milligrams of TNT were placed atop 10 cm of snow-March through 27 March 1989. Snow samples were collected from the contemporary surface 15 cm from the specimen, on 14 and 24 March, and from the surface 8 cm above the specimen, in new snow 21 March. Analysis of the specimens by D. Leggett showed presence of the tracer, with a variation in concentration of about a factor of two over these distances. This, or a similar organic material may provide a useful tracer for future studies.

CONCLUSIONS:

Sparcity of snowfall and snowpack at Whiteface prevented completion of a compreprecipitation experiment during winter 1988-89. An instrumented snow field was established, and the instrumentation and recording system produced data with sufficient accuracy and time resolution to support the planned experiment. An organic tracer produced a detectable signal after diffusing through the

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FIG I. TEMP 3cm BELOW GRND/AIR INTERFACE

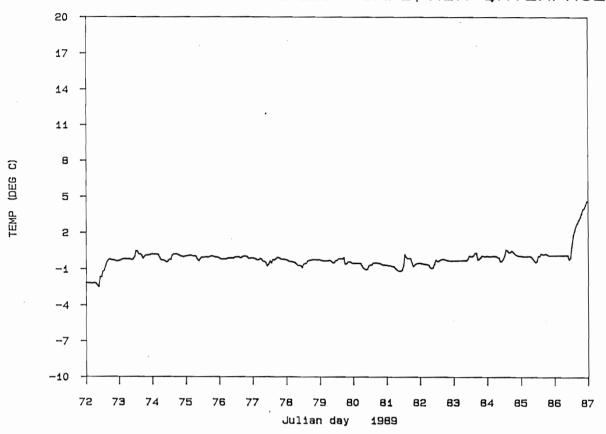
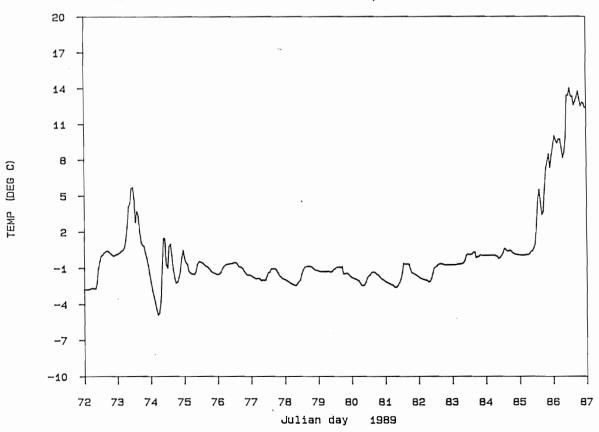
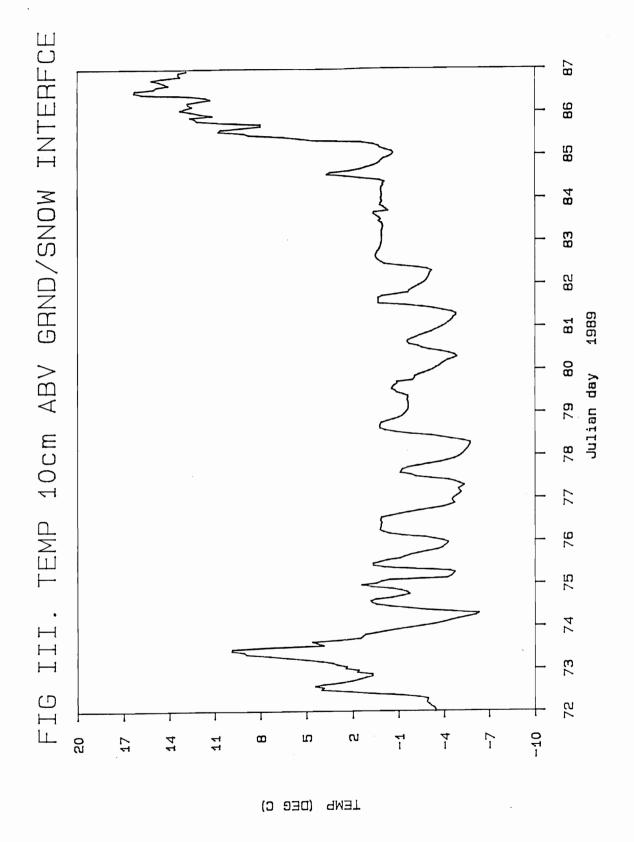


FIG II. TEMP AT GRND/SNOW INTERFACE

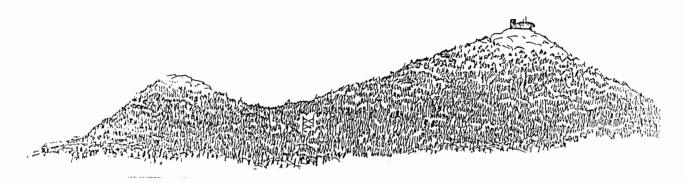




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Base (600m)

Ozone
Meteorological
wind speed
wind direction
temperature
solar radiation
relative humidity
barometricpressure
Filter (tripack)
aerosols
nitric acid
sulfur dioxide

Dry deposition (NOAA/ADTL) Precipitation as: wet-only (MAPS3S,NADP)

Condensation nuclei (summer)

Tower (1000m)

Ozone
Meteorological
wind speed
wind direction
temperature
solar radiation
relative humidity

Filter (tripack)
aerosols
nitric acid
sulfur dioxide

Dry deposition

Precipitation as: wet-only

cloudwater
rime ice
throughfall
stemflow

Summit (1483m)

Ozone
Meteorological
wind speed
wind direction
temperature

barometric pressure
Filter (high-volume)
aerosols
nitric acid
sulfur dioxide
Gaseous
sulfur dioxide
nitrous oxides

Precipitation as:

cloudwater rime ice

Condensation nuclei (summer) Hydrogen Peroxide

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